Egyptian Geoid using Ultra High-Degree Tailored Geopotential Model

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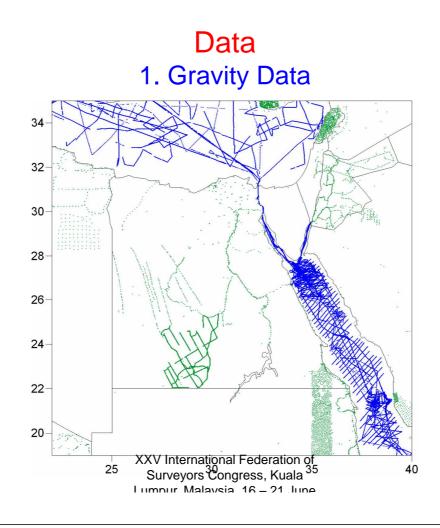


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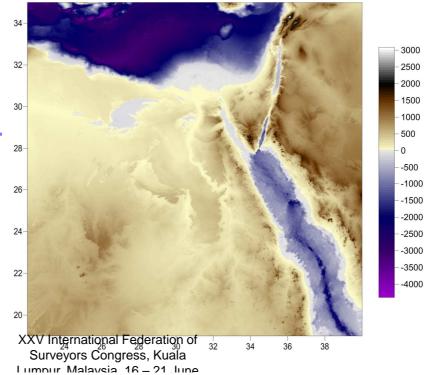
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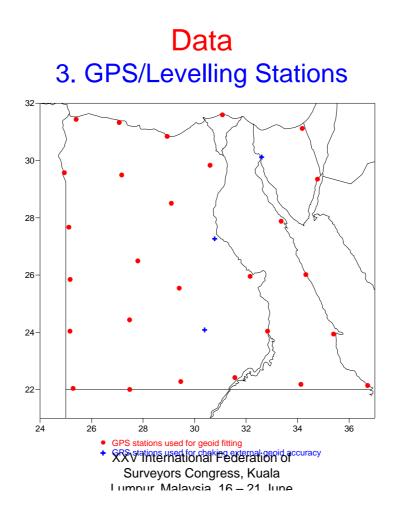


Data 2. Digital Height Models (EGH13) (Abd-Elmotaal and Ashry, 2013)

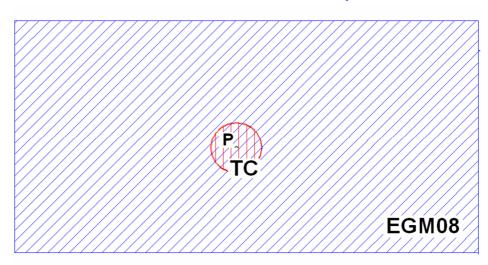


-4000

- ♣ Coarse DHM resolution is 30" × 30"
- ♣ Fine DHM resolution is 3" × 3"



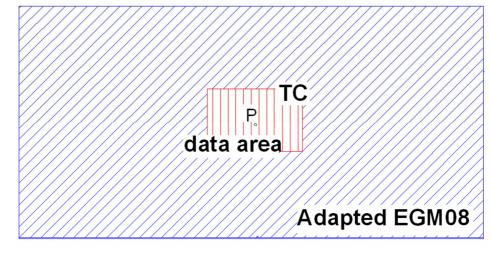
Remove/Restore Technique 1. Traditional Technique



Double consideration of the topographic-isostatic masses within the radius R₂ (double hatched area)

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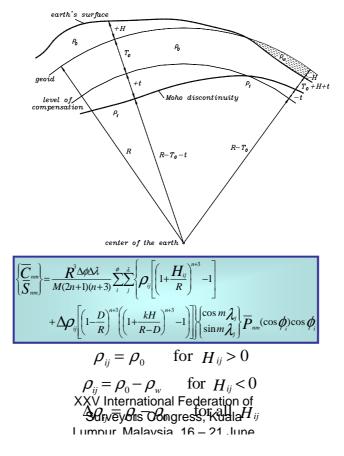
Remove/Restore Technique 2. Window Technique



No double consideration of the topographic-isostatic masses (no double hatching)

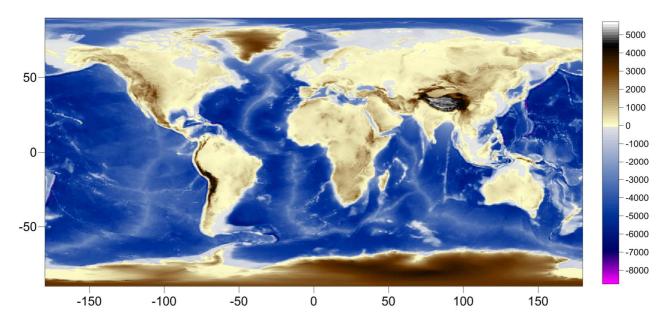
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Harmonic Analysis of the Topographic-Isostatic Potential



Methodology

Preparing Gravity Anomalies for Tailored Geopotential Model Computation 1.1. SRTM 30' × 30' Global Digital Height Model



Min = -8756 m

Max = 5719 m Mean = -1892 m XXV International Federation of Surveyors Congress, Kuala St. dev. = 2638 m

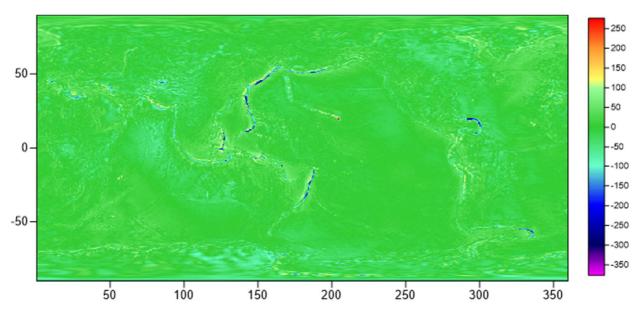
Preparing Gravity Anomalies for Tailored Geopotential Model Computation

1.2. EGM2008TI Topographic-Isostatic Geopotential Model

$$\left\{ \overline{C}_{nm} \\ \overline{S}_{nm} \right\}_{EGM \ 2008TI} = \left\{ \overline{C}_{nm} \\ \overline{S}_{nm} \right\}_{EGM \ 2008} - \left\{ \overline{C}_{nm} \\ \overline{S}_{nm} \right\}_{TI}$$

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Preparing Gravity Anomalies for Tailored Geopotential Model Computation 2. Global isostatic Anomalies (EGM2008TI)



Min = -337 mgal

Max = 273 mgal Mean = -1 mgal XXV International Federation of Surveyors Congress, Kuala St. dev. = 27 mgal

Preparing Gravity Anomalies for Tailored Geopotential Model Computation 3. Local Isostatic (Free-air - window topo-iso) Gravity Anomalies (Egypt) 34 220 32 200 180 160 2,0 140 30 120 100 80 60 28 40 20 0 -20 26 -40 -60 -80 24 -100 -120 -140 -160

Preparing Gravity Anomalies for Tailored Geopotential Model Computation 4. Merged Isostatic Gravity Anomalies (EGM2008TI + Egypt)

30

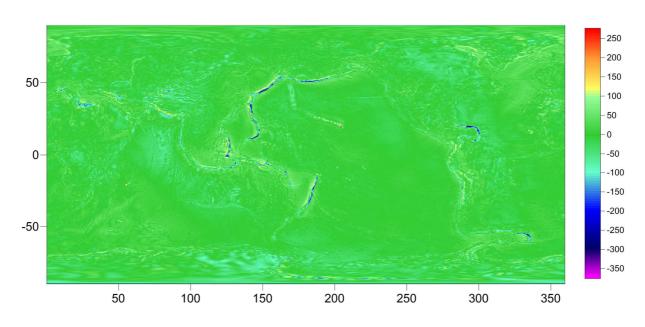
32

Max = 202 Sugelyors Congless, Kutala mgal

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34

36



Min = -337 mgal

22

20

Min = -190 mgal

Max = 273 mgal Mean = -1 mgal XXV International Federation of Surveyors Congress, Kuala St. dev. = 27 mgal

-180 -200

St. dev. = 37 mgal

Harmonic Analysis using FFT Technique

4 Expanding $f(\theta, \lambda) \rightarrow$

$$f(\theta,\lambda) = \sum_{n=0}^{\infty} \sum_{m=-n}^{n} \overline{F}_{nm} \, \overline{Y}_{nm}(\theta,\lambda)$$

♣ Orthogonality →

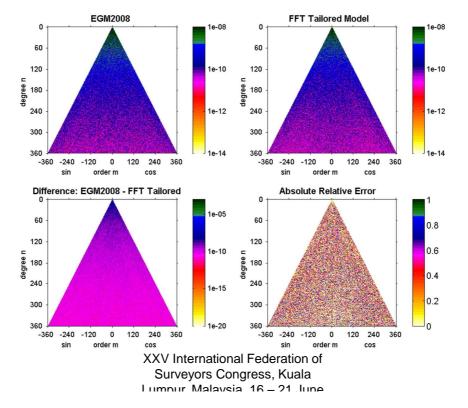
$$\overline{F}_{nm} = \frac{1}{4\pi} \iint_{\sigma} f(\theta, \lambda) \overline{Y}_{nm}(\theta, \lambda) \, d\sigma$$

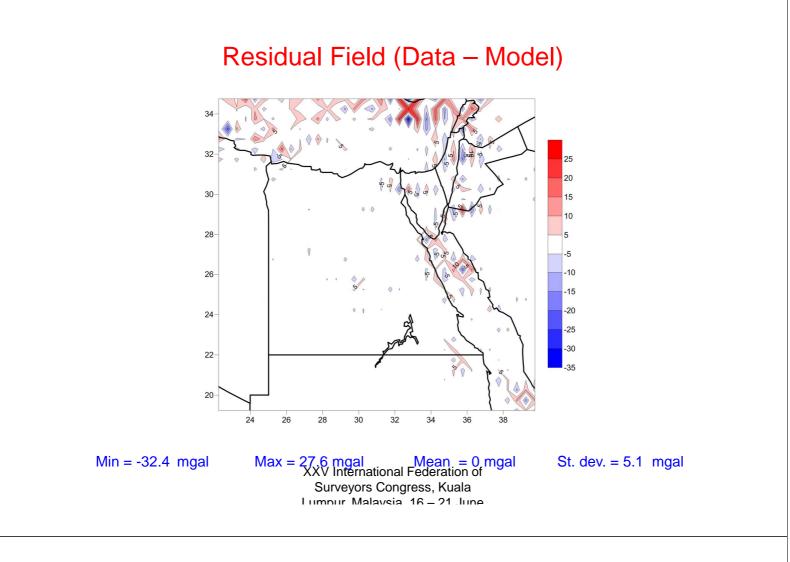
- HRCOFITR Program (Abd-Elmotaal, 2004 b) has been used
- Iteration to obtain the best coefficients accuracy and minimum residual field

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Tailored Reference Geopotential Model for Egypt (EGTGM2014)

(lower harmonics only till n=360; model is available till n=2160)

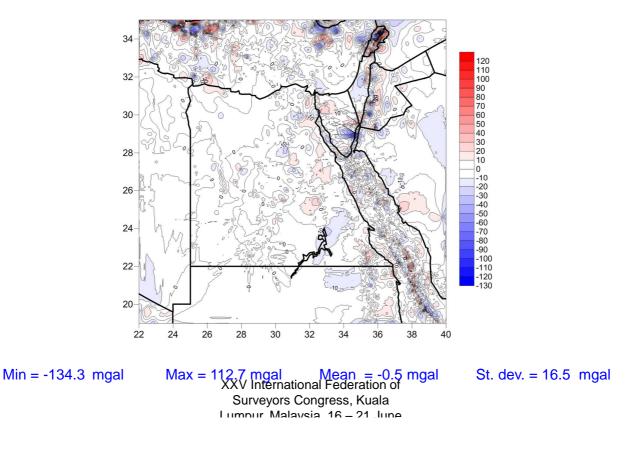




Gravity Reduction (102418 points)

Reduced gravity	Minimum	Maximum	Average	St. dev.	Variance
Δg_{F}	-210.60	314.99	-27.58	50.65	2565.1
$\Delta g_{Airy\ window}$ (EGM2008)	-99.15	122.49	-0.26	20.46	418.6
$\Delta g_{Airy window}$ (EGTGM2014)	-134.30	112.69	-0.45	16.45	270.5

Airy window isostatic anomalies for Egypt using the EGTGM2014 tailored geopotential model



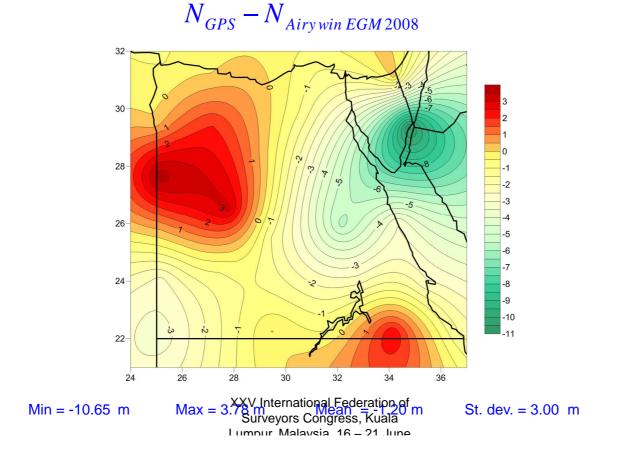
Stokes' Integral with 1D-FFT Technique

$$N = \frac{R}{4\pi\gamma} \iint_{\sigma} \Delta g \ S(\psi) \ d\sigma$$
$$N(\phi_{P}, \lambda_{P}) = \frac{R \Delta \phi \Delta \lambda}{4\pi\gamma} \sum_{Q} (\Delta g_{Q} \cos \phi_{Q}) S(\psi_{PQ})$$
$$N_{\phi_{P}}(\lambda) = \frac{R \Delta \phi \Delta \lambda}{4\pi\gamma} \sum_{\phi_{Q}} \cos \phi_{Q} \sum_{\lambda_{Q}} \Delta g_{Q} \ S(\psi_{PQ})$$
$$S(\psi_{PQ}) = S_{\phi_{P}\phi_{Q}}(\lambda_{P} - \lambda_{Q}) = S_{\phi}(\lambda_{P} - \lambda_{Q})$$
$$N_{\phi_{P}}(\lambda) = \frac{R \Delta \phi \Delta \lambda}{4\pi\gamma} \sum_{\phi_{Q}} \cos \phi_{Q} \sum_{\lambda_{Q}} \Delta g_{Q} \ S_{\phi}(\lambda_{P} - \lambda_{Q})$$

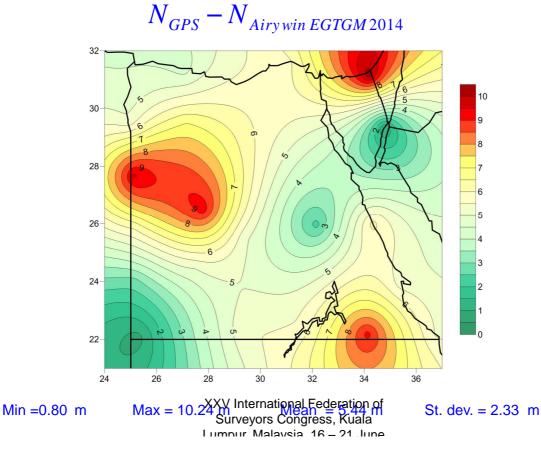
$$N_{\phi_{P}}(\lambda) = \frac{R \Delta \phi \Delta \lambda}{4\pi \gamma} \mathfrak{P}_{1}^{-1} \left\{ \sum_{\phi_{Q}} \mathfrak{P}_{1}(\Delta g_{Q} \cos \phi_{Q}) \mathfrak{P}_{1}[S_{\phi}(\lambda_{P} - \lambda_{Q})] \right\}$$

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Absolute Geoid Difference to GPS/levelling



Absolute Geoid Difference to GPS/levelling



Statistics of the Remaining Differences at the 27 GPS Stations used for the Geoid Fitting after Removing a Kriging Trend Function (Internal Precision)

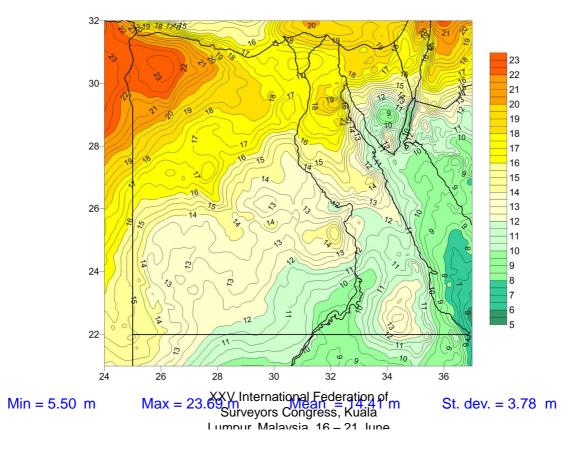
Geoid type	Minimum [cm]	Maximum [cm]	Average [cm]	Standard deviation [cm]
$N_{AirywinEGM2008}$	-9.2	9.8	0.2	3.4
$N_{Airywin EGTGM 2014}$	-7.6	9.8	0.2	3.3

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Statistics of the Remaining Differences after Removing a Kriging Trend Function at the 3 GPS Stations which were not used for the Geoid Fitting (External Accuracy)

Geoid type	Minimum [m]	Maximum [m]	Average [m]	Standard deviation [m]
$N_{AirywinEGM2008}$	-1.40	2.46	0.29	1.97
$N_{AirywinEGTGM2014}$	-0.88	2.13	0.30	1.60

Gravimetric Geoid for Egypt using EGTGM2014



Summary and Conclusions

- An ultra high-degree tailored reference geopotential model for Egypt, complete to degree and order 2160, has been developed.
- The tailored geopotential model created in this investigation gives better residual gravity anomalies (unbiased and have much less variance). The variance has dropped by about 35%.
- Gravimetric geoids for Egypt have been computed in this investigation using both the EGM2088 and the EGTGM2014 tailored geopotential models in the framework of the window remove-restore technique using 1D-FFT technique.
- The computed geoids have been fitted to the GPS-derived geoid by removing a trend surface.
- A kriging trend function has been computed using only 27 GPS stations among the available 30 GPS stations in Egypt.
- The internal precision of the fitted geoids is very good (about 3 cm) and it is nearly equal for both geoids.
- Using the EGTGM2014 tailored geopotential model improves the external geoid accuracy by about 20%, and the range of the remaining differences has dropped by about 22%.

Thank You!

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