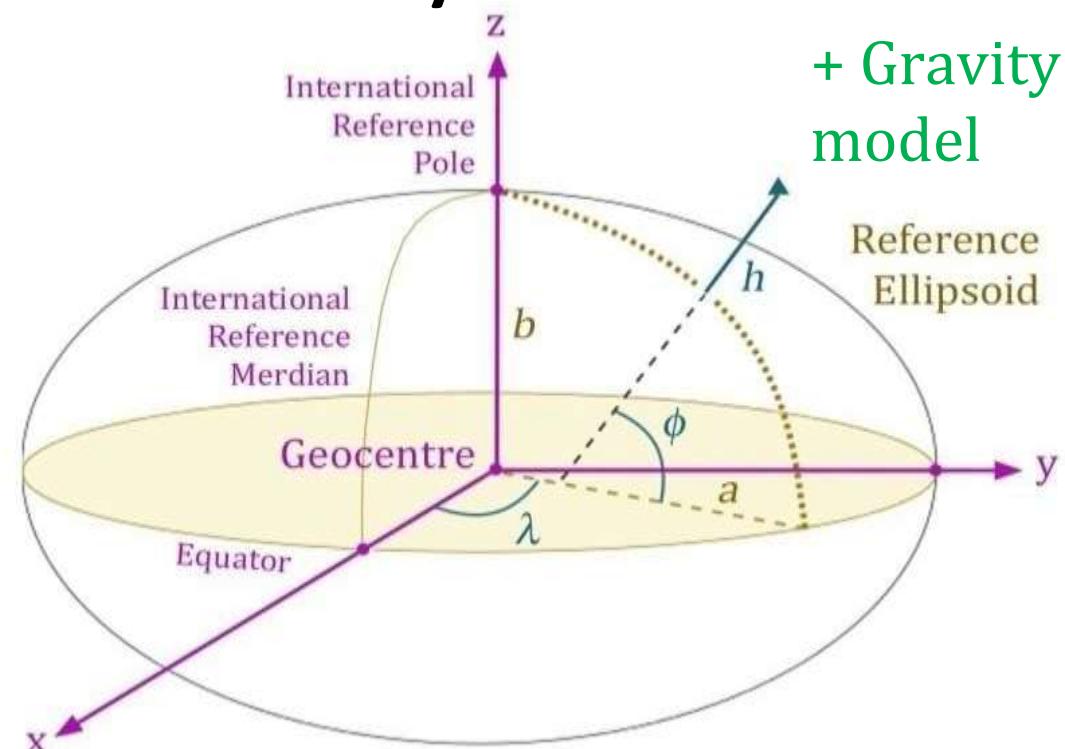


# Reference Frames, Transformations and GIS

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The Ohio State University  
USA

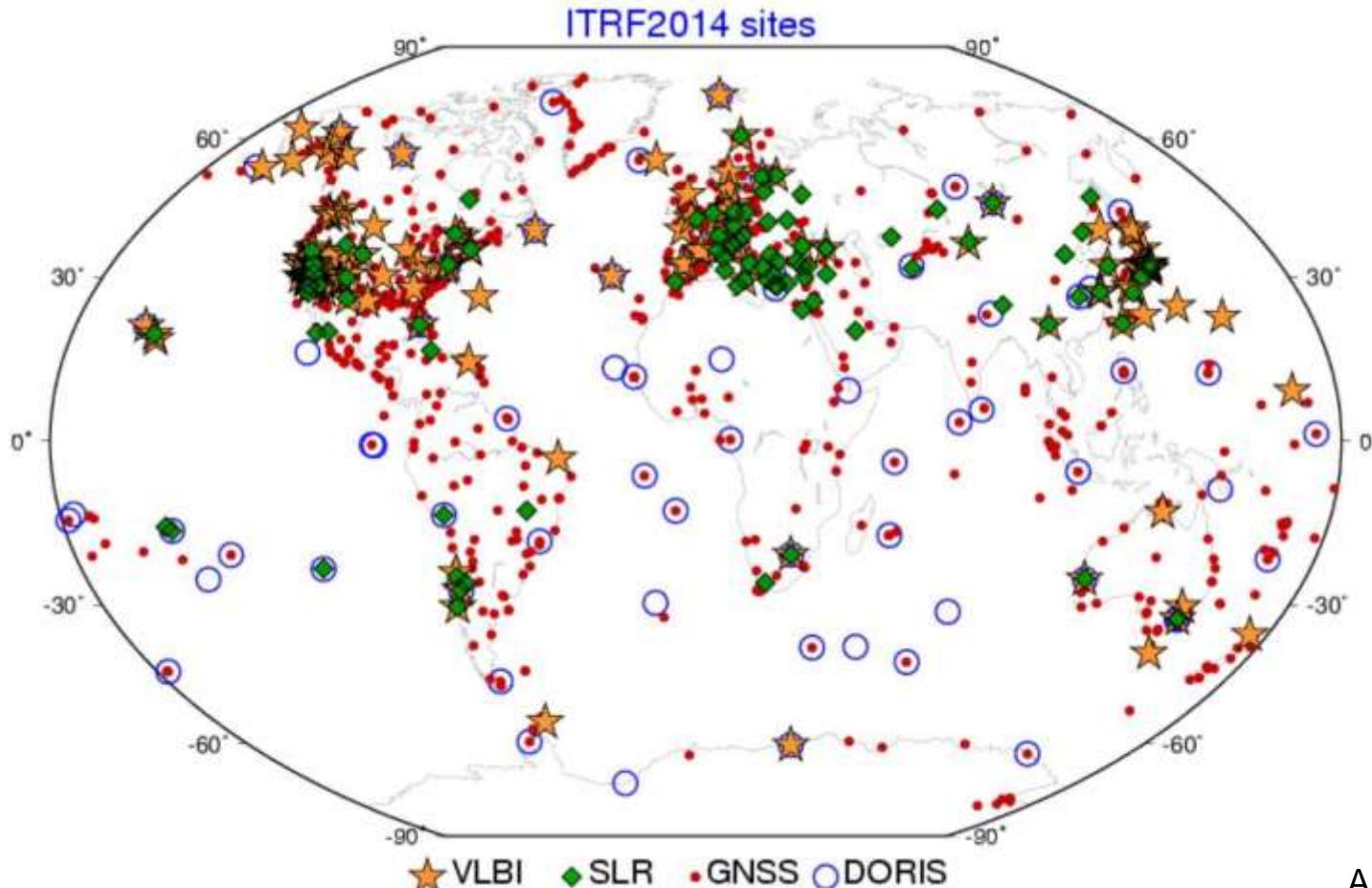
# Terrestrial Reference System

In principle a TRS  
should be  
invariant with time



example: Geocentric ITRS

### International Terrestrial Reference Frame

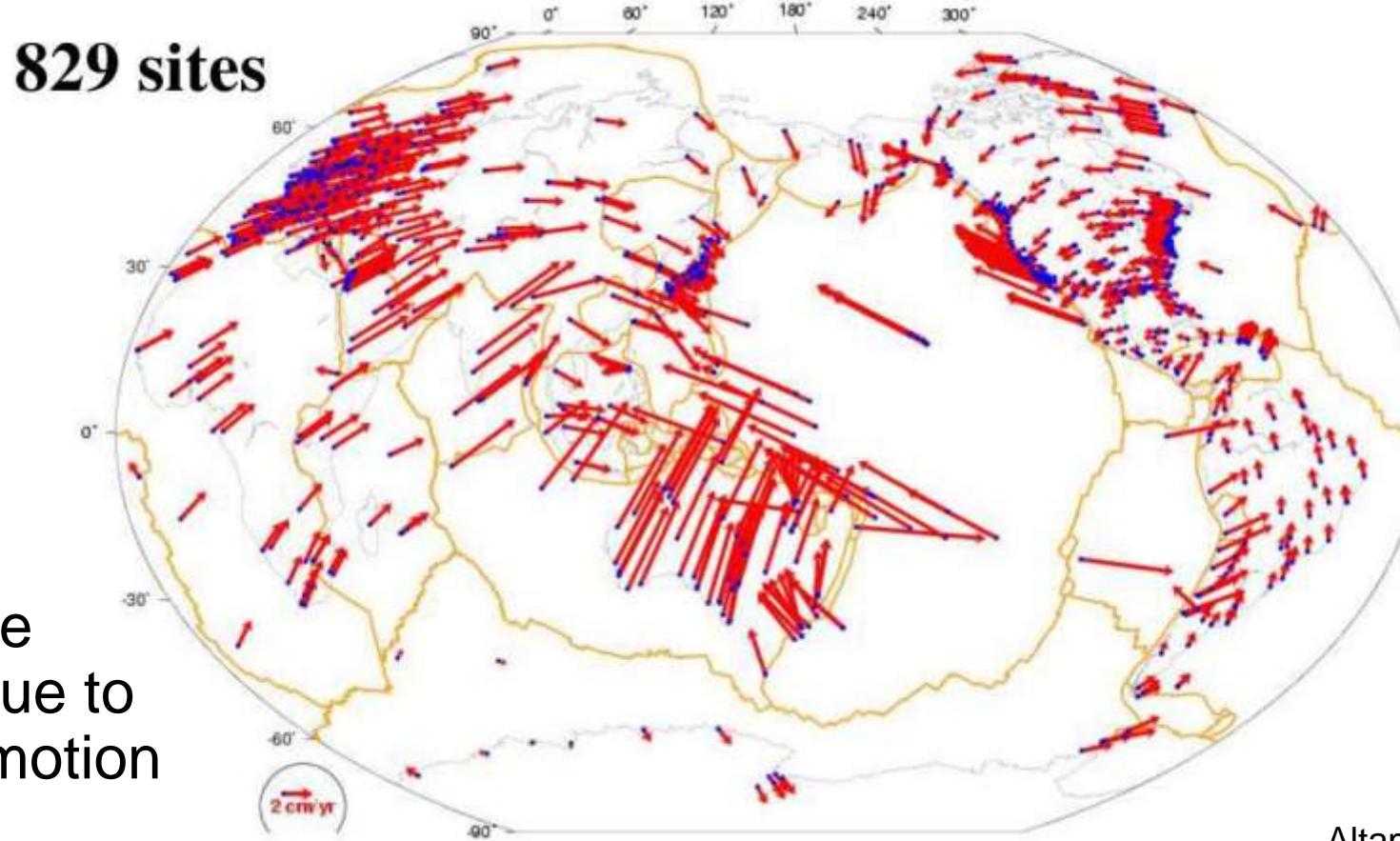


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## Reference Frames in Practice



ITRF Kinematics – NNR site velocities (the angular momentum of all tectonic plates sums cero)



Altamimi et al., 2015

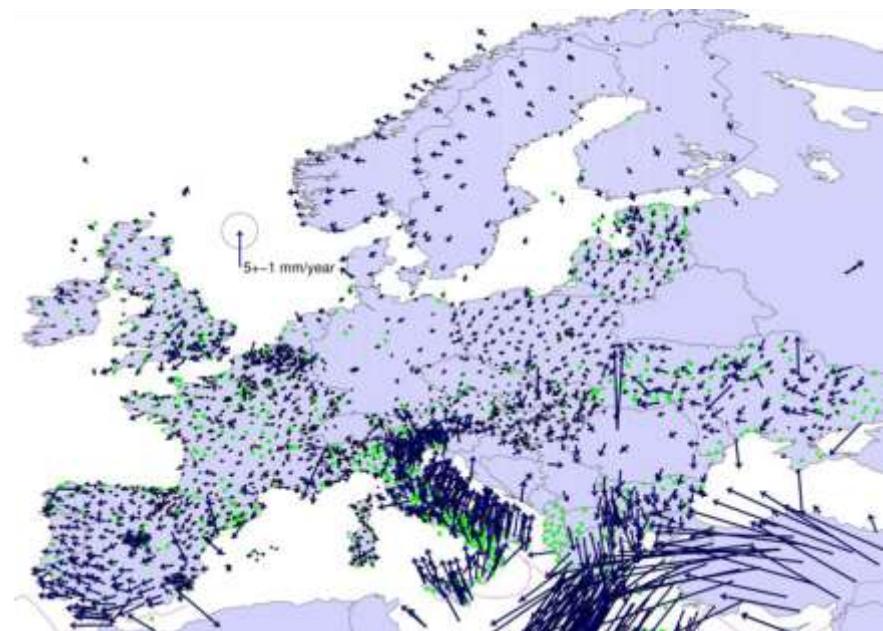
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## Reference Frames in Practice



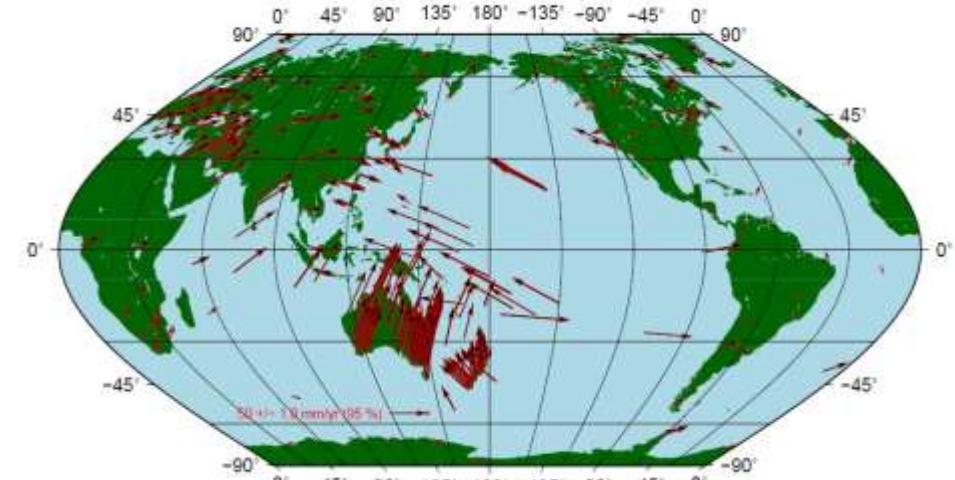
### Regional Reference Frames

Plate fixed

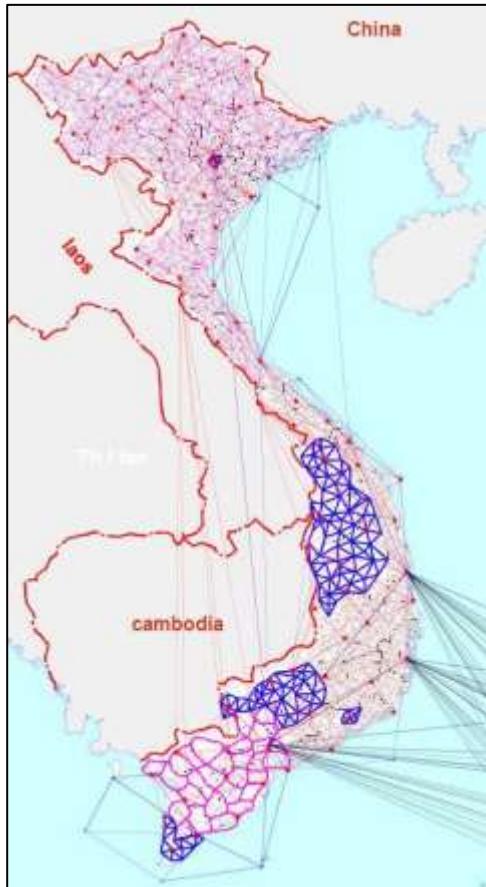


e.g. European RF (ETRF) – velocities minimised for most of Eurasian plate  
(figure: EUREF, 2018)

NNR (No-Net-Rotation)



e.g. Asia-Pacific RF (APREF) – no dominant plate within frame coverage so NNR model is used (figure: Hu, 2014)



## National & Local Reference Frames

Characteristics:

Now usually a fixed epoch of ITRF using GNSS (recent geocentric frames) e.g. VN2000 for mainland Vietnam (figure: Vietnam Dept. of Surveying Mapping, 2016)

or astronomical determination of frame origin (pre space geodetic era)  
(frame often not geocentric)

e.g. Australian Geodetic Datum 1966 (AGD66)

(figure: Paul Wise, 2016)



## Static and time-dependent frames

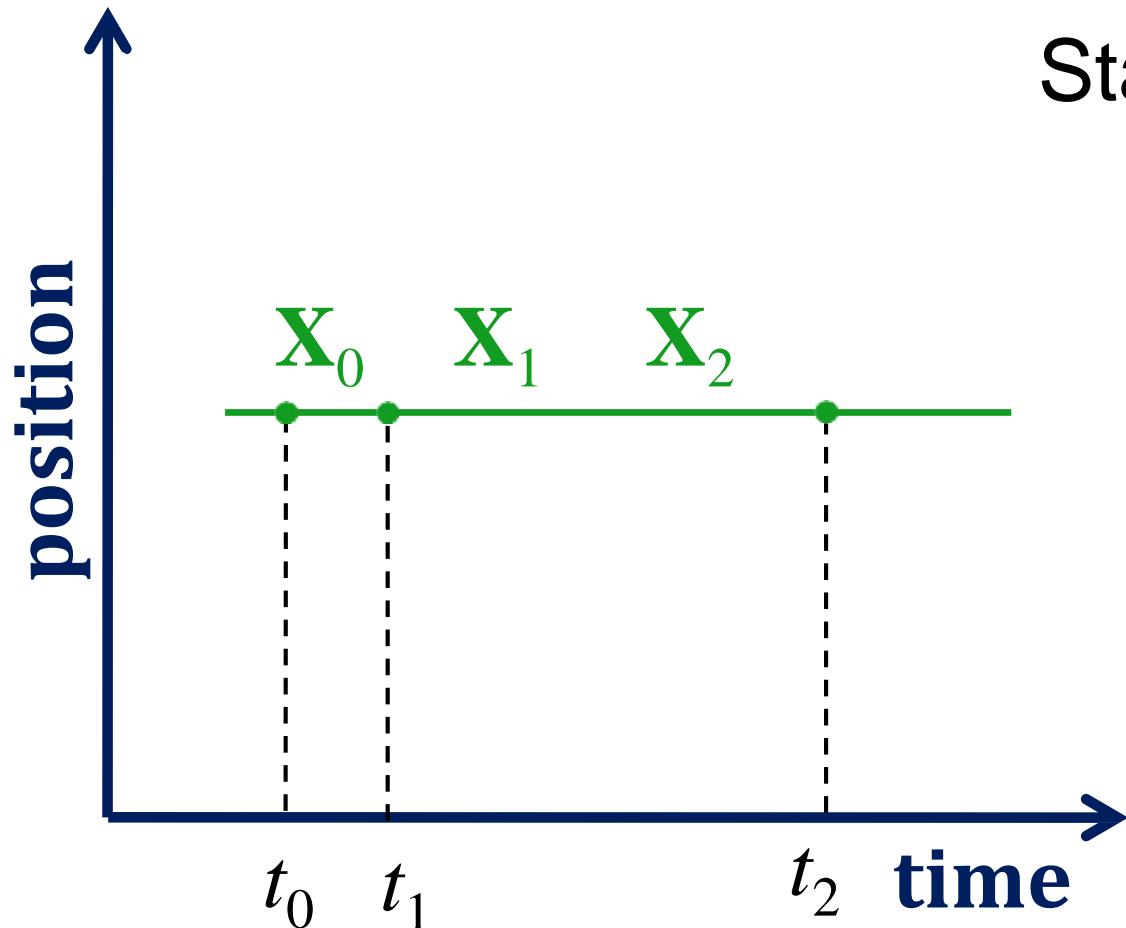
Static (time-invariant) – no displacement is assumed

Kinematic (time dependent)

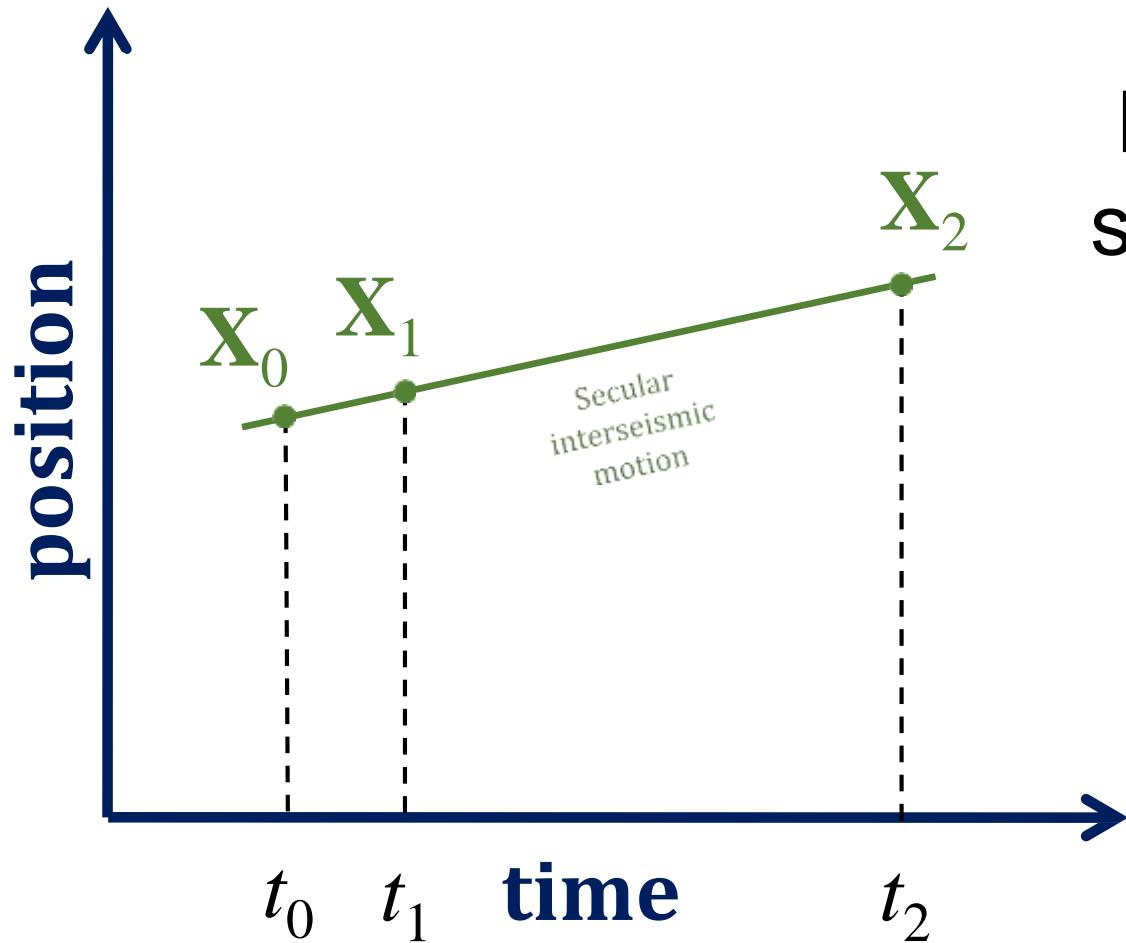
- includes all displacements wrt. ITRS  
(e.g., tectonic motion, glacial isostatic adjustment, coseismic deformation)

Semi-kinematic (time dependent)

- secular tectonic motion is modelled out
- coseismic displacement models applied after earthquakes

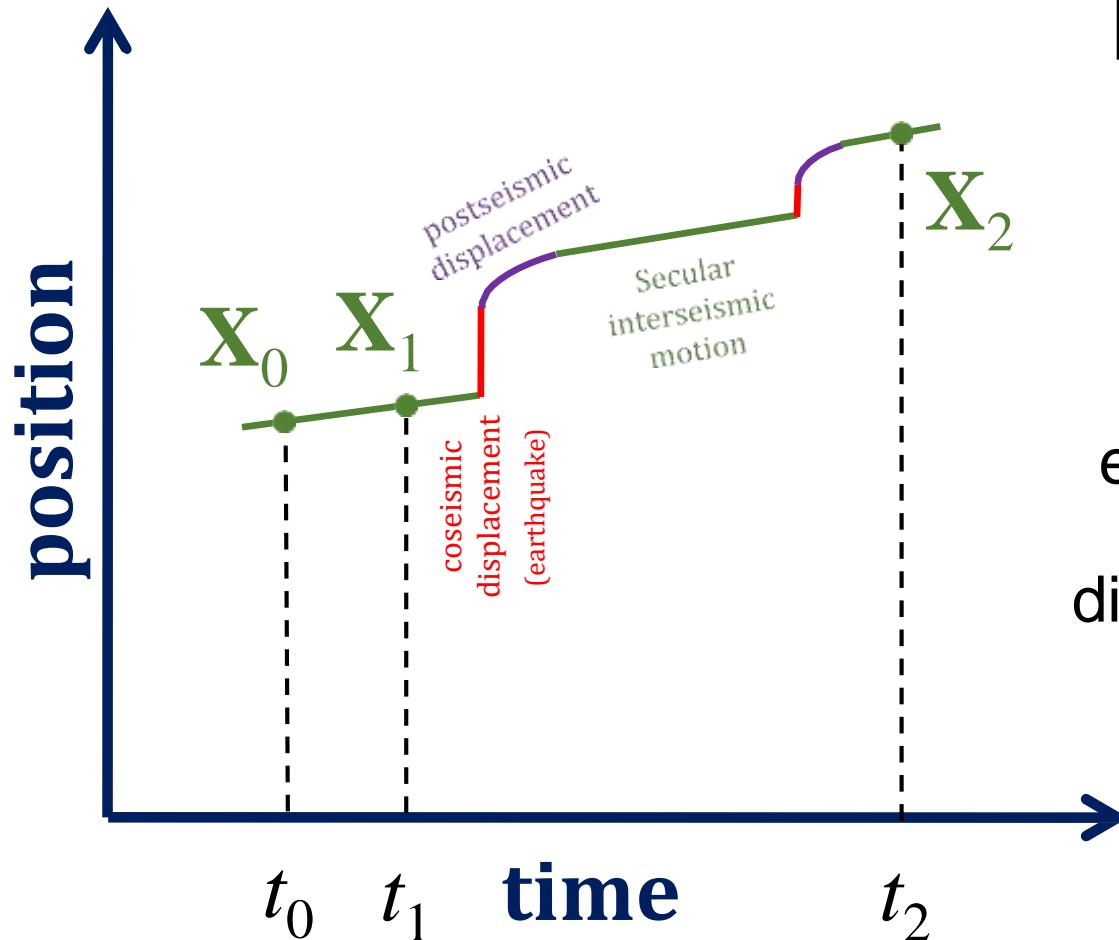


Static reference frame  
position does not  
change with time  
(except perhaps for  
readjustments)



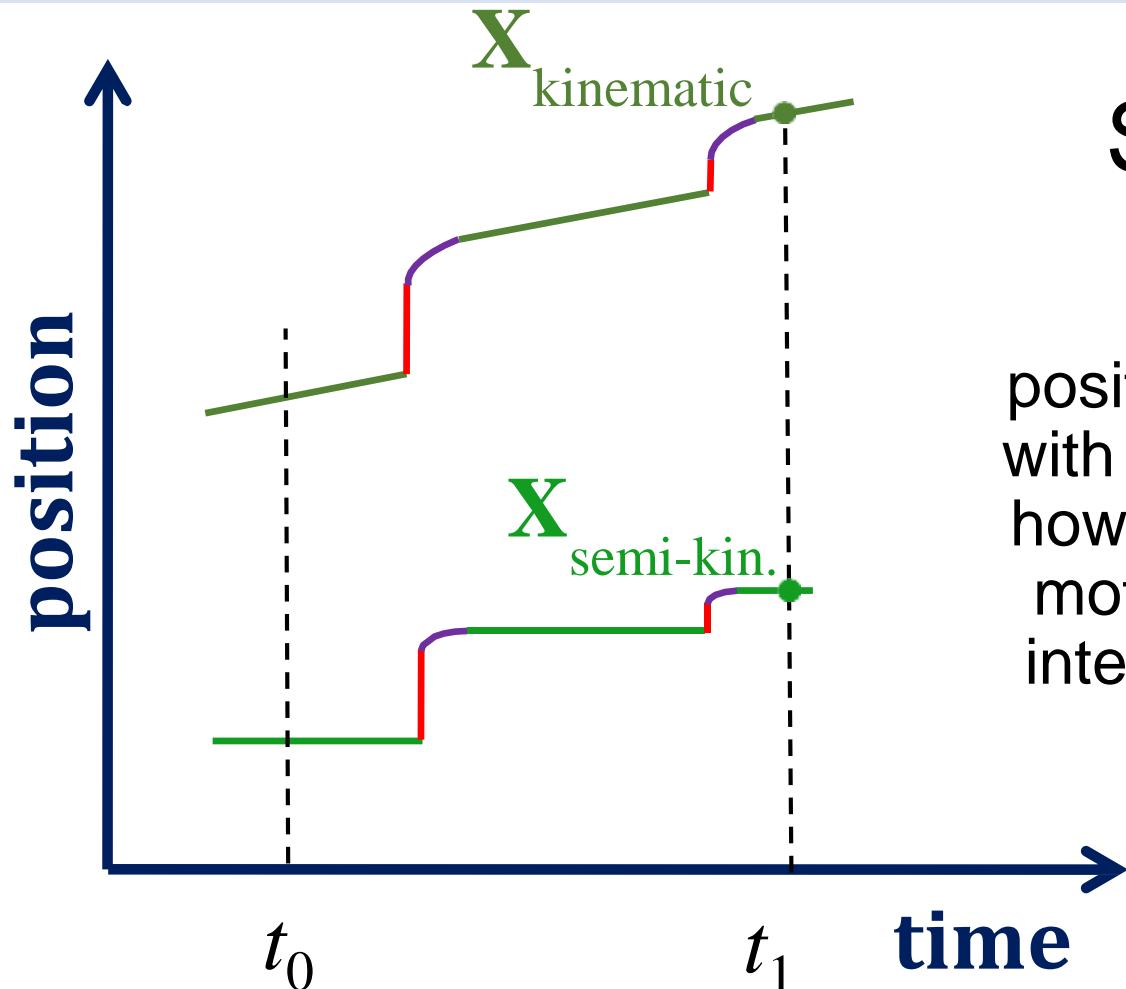
Kinematic frame –  
stable plate regions

position changes  
slowly ( $< 8 \text{ cm yr}^{-1}$ )  
due to secular plate  
motion and GIA  
(interseismic motion)



Kinematic frame – plate boundary zones

position changes episodically with coseismic and postseismic displacements resulting from earthquakes



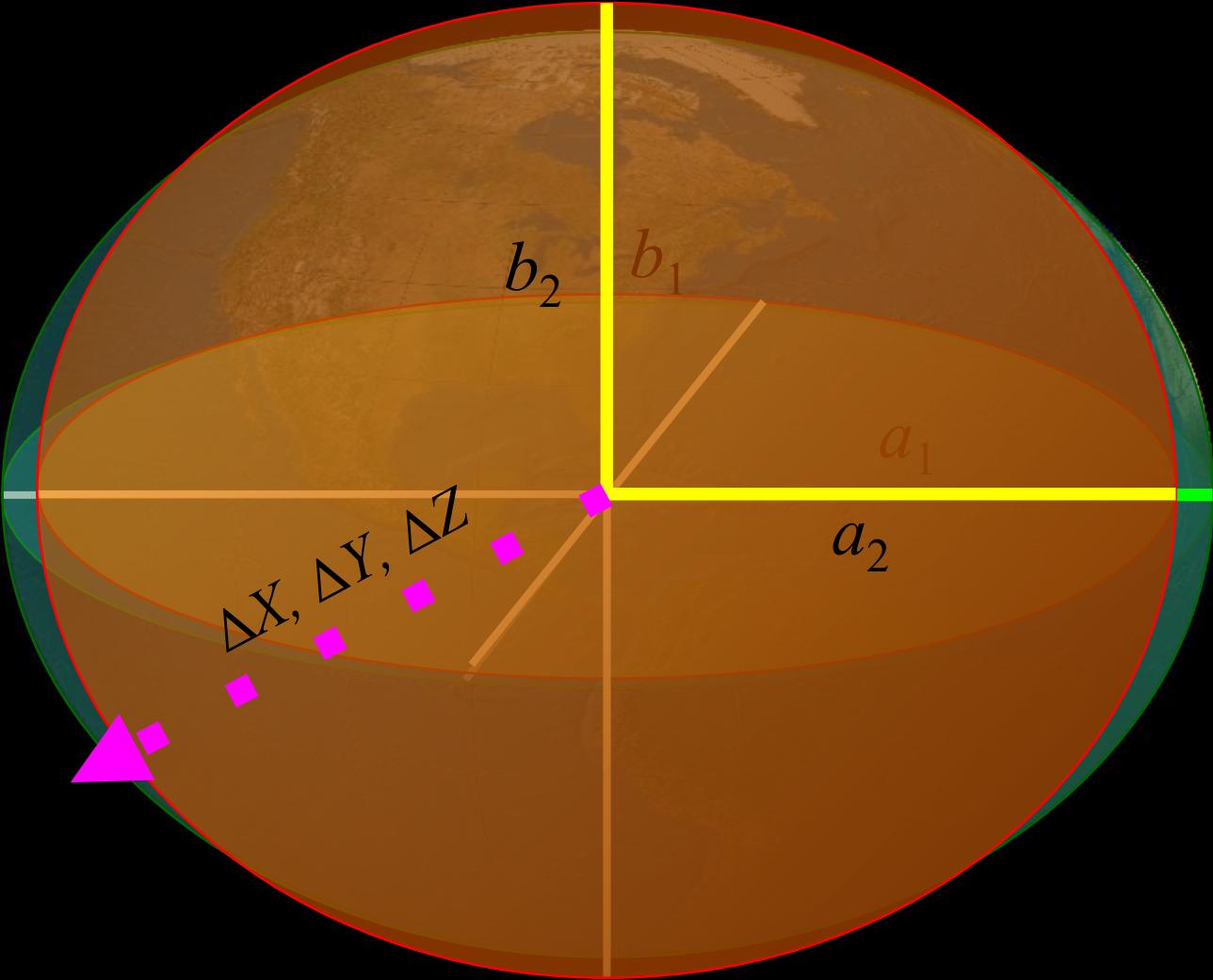
## Semi-kinematic frame

position changes episodically with coseismic displacements however secular interseismic motion is modelled out until interseismic strain threshold is reached

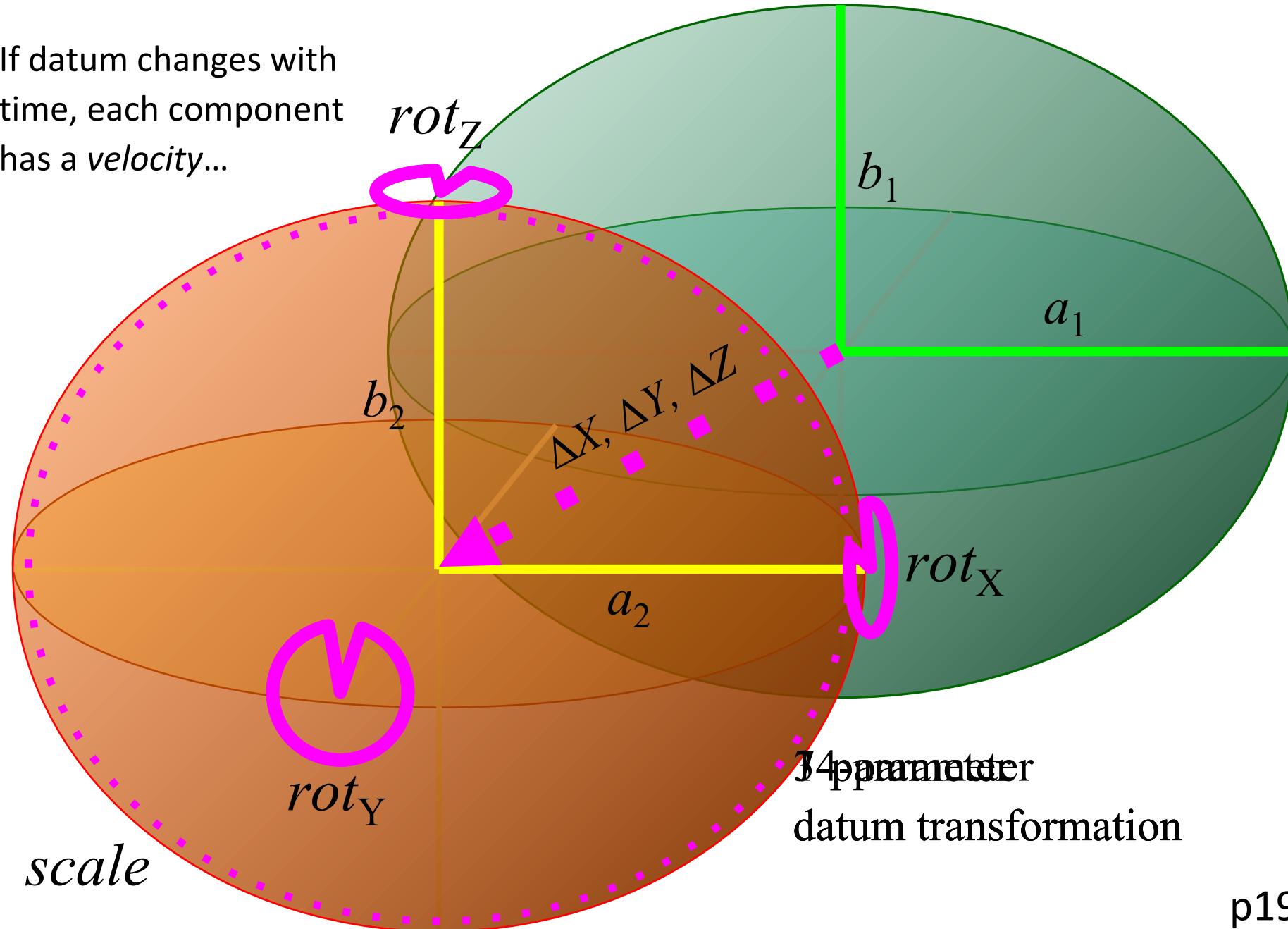
## Interframe transformations

	Static frame	time-dependent frame
Conformal Parametric	3 parameter  7 parameter	14 parameter (7 parameters at reference epoch + rates)
Grid model	displacement grid  distortion grid	interseismic velocity model + coseismic displacement grid + postseismic decay term grid

# Geometric datum transformations



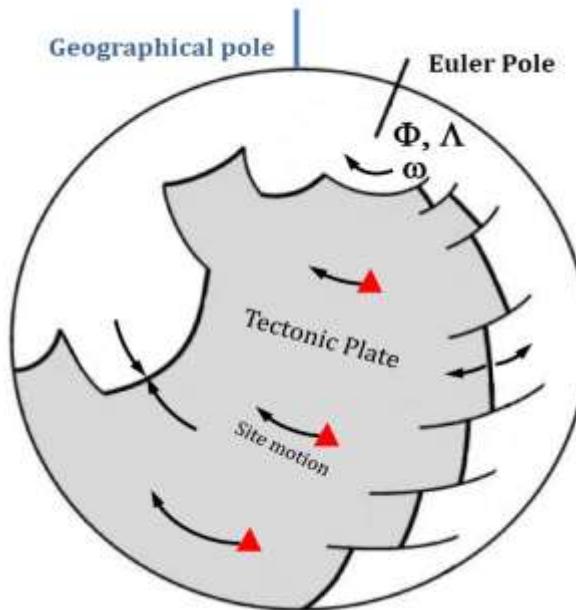
If datum changes with time, each component has a *velocity*...



# Datum transformations

- Typical geodetic datum transformations
  - **3-parameter:** 3-dimensional translation of origin as  $\Delta X$ ,  $\Delta Y$ ,  $\Delta Z$  (*just like a GPS vector*)
  - **7-parameter:** 3 translations *plus* 3 rotations (one about each of the axes) *plus* a scale
  - **14-parameter:** A 7-parameter where each parameter changes with time (each has a *velocity*)
  - Transformations that also model distortion and tectonic motion, such as the NGS programs NCCAT and HTDP
- Vertical datum transformations
  - Can be simple shift or complex operations that model distortion, gravity, and tides, such as NGS programs NCCAT and VDatum

# Plate Motion Model



$$\begin{aligned}\Omega_x &= \cos(\Phi) \\ \Omega_y &= -\sin(\Phi) \cos(\Lambda) \omega \\ \Omega_z &= \sin(\Phi) \omega\end{aligned}$$

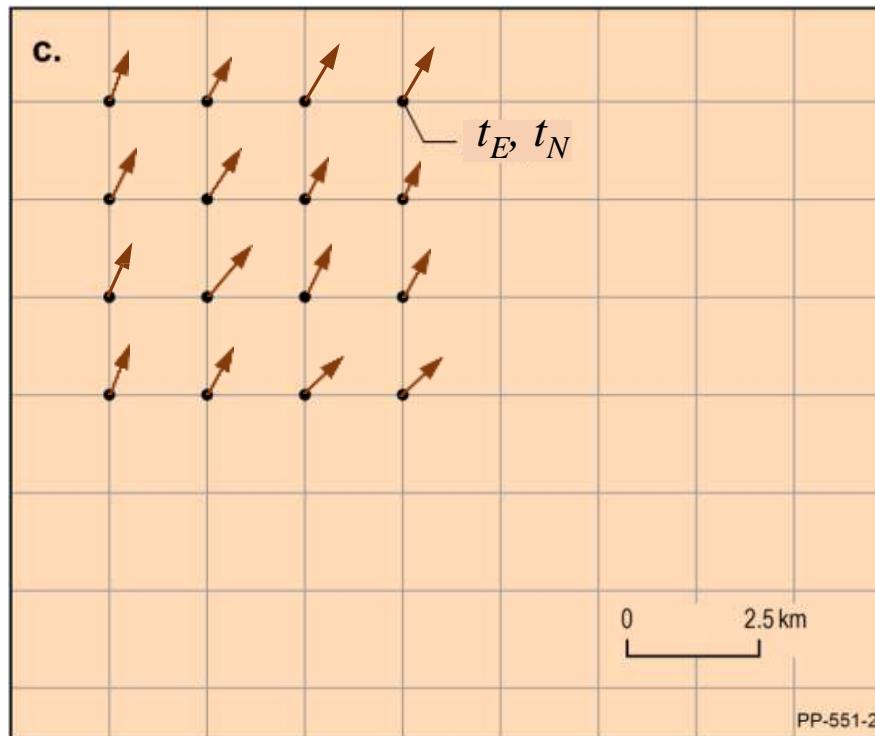
Plate	Euler pole of rotation			Equivalent Cartesian angular velocity		
	$\Phi$ (°)	$\Lambda$ (°)	$\omega$ (° Ma <sup>-1</sup> )	$\Omega_x$ (Rad Ma <sup>-1</sup> )	$\Omega_y$ (Rad Ma <sup>-1</sup> )	$\Omega_z$ (Rad Ma <sup>-1</sup> )
Antarctic	58.8	-127.4	0.219	-0.001202	-0.001571	0.003272
Arabian	51.2	-6.7	0.515	0.005595	-0.000659	0.007001
Australian	32.4	38.1	0.631	0.007321	0.005730	0.005890
Eurasian	55.1	-99.1	0.261	-0.000412	-0.002574	0.003733
Indian	51.6	-0.2	0.516	0.005595	-0.000024	0.007049
Nazca	45.8	-102.2	0.629	-0.001614	-0.007486	0.007869
North American	-5.2	-88.0	0.194	0.000116	-0.003365	-0.000305
Nubian	49.7	-80.8	0.267	0.000480	-0.002977	0.003554
Pacific	-62.6	111.3	0.679	-0.001983	0.005076	-0.010516
South American	-19.1	-131.9	0.119	-0.001309	-0.001459	-0.000679
Somalian	47.7	-98.7	0.332	-0.000587	-0.003849	0.004286

ITRF2014 Plate Motion Model (from Altamimi *et al.*, 2017)

Euler pole can be estimated by LS inversion of interseismic site velocities within a stable plate

Rotation rates can be used in a 14-parameter model with zeroes for other parameters (now used for GDA2020 to ITRF2014 transformations)

# Displacement and Distortion Grids



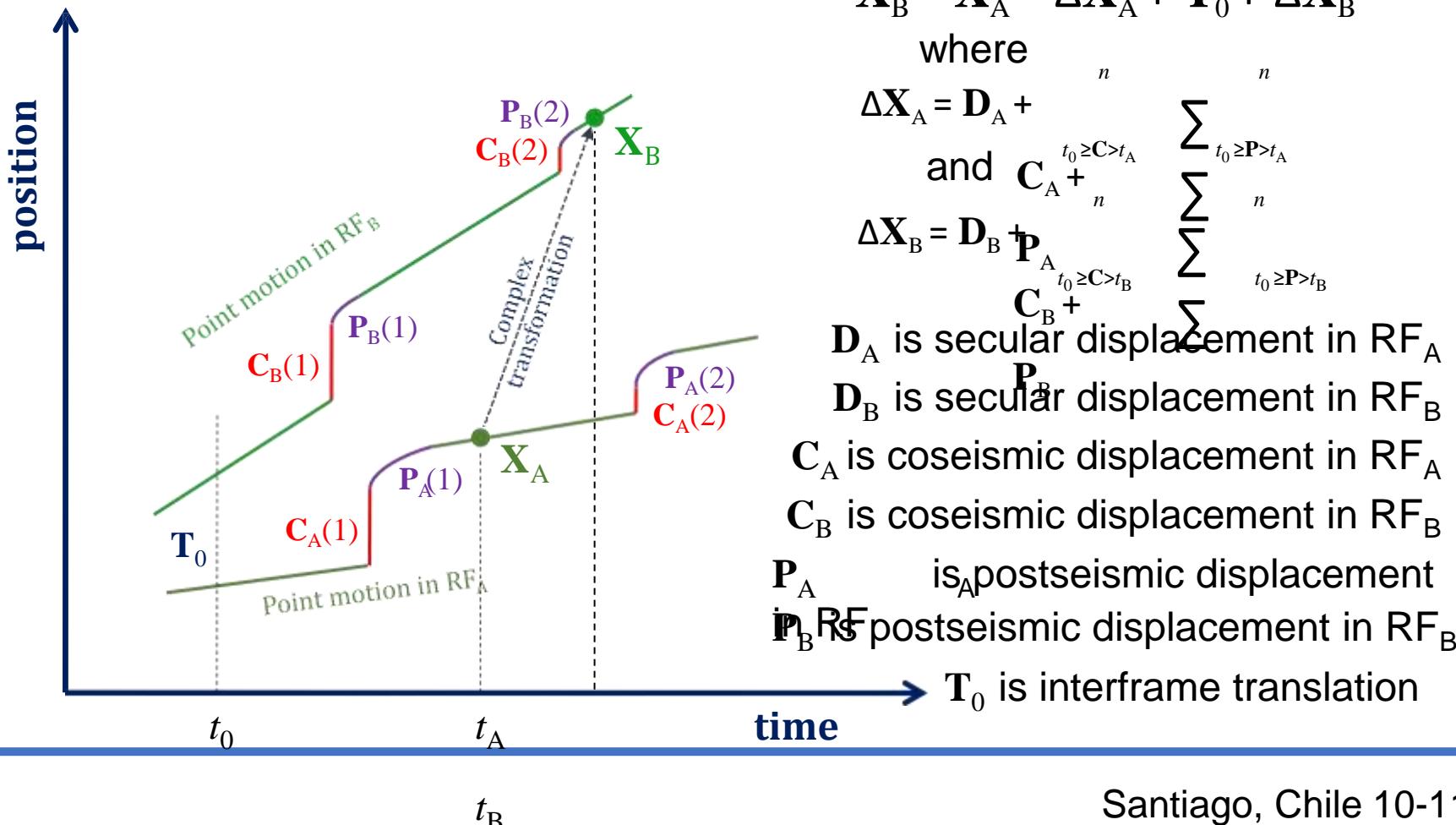
adapted from GDA2020 Technical Manual, ICSM,  
2018

$$\begin{bmatrix} E \\ N \end{bmatrix} = \begin{bmatrix} E \\ N \end{bmatrix}_A + \begin{bmatrix} t_E \\ t_N \end{bmatrix}$$

Topocentric shifts estimated by  
bilinear interpolation of grid model  
(e.g., in NTv2 format)

Ideal for heterogeneous RF  
transformations and handling local  
distortions

## Complex time-dependent transformations



# Topocentric frame

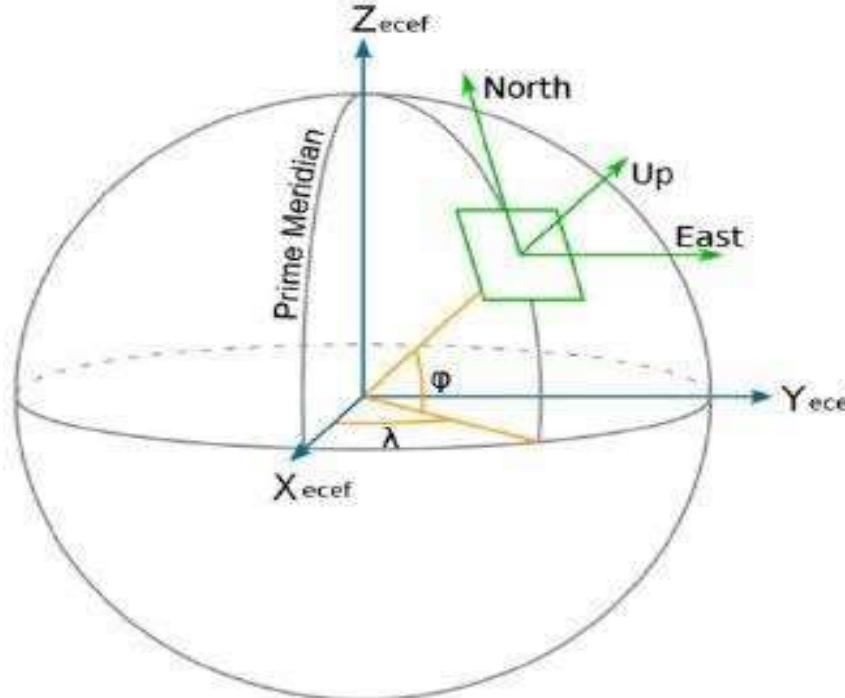


Image: Haasdyk and Janssen, 2011

A topocentric projection canvas is useful for complex transformation computations involving displacement grids

Converting geocentric translation to topocentric (rate or shift)

$$\begin{bmatrix} 1 \\ -\sin\lambda & \cos\lambda & 0 \\ U & \sin\varphi & 0 \end{bmatrix}$$

Converting topocentric translation to geocentric (rate or shift)

$$\begin{bmatrix} X \\ E \\ Z \end{bmatrix} = \begin{bmatrix} 1 & -\sin\lambda & \cos\lambda \\ 0 & -\sin\varphi\cos\lambda & \cos\varphi\cos\lambda \\ 0 & \sin\varphi & 0 \end{bmatrix} \begin{bmatrix} U \\ V \\ W \end{bmatrix}$$

# GIS transformations

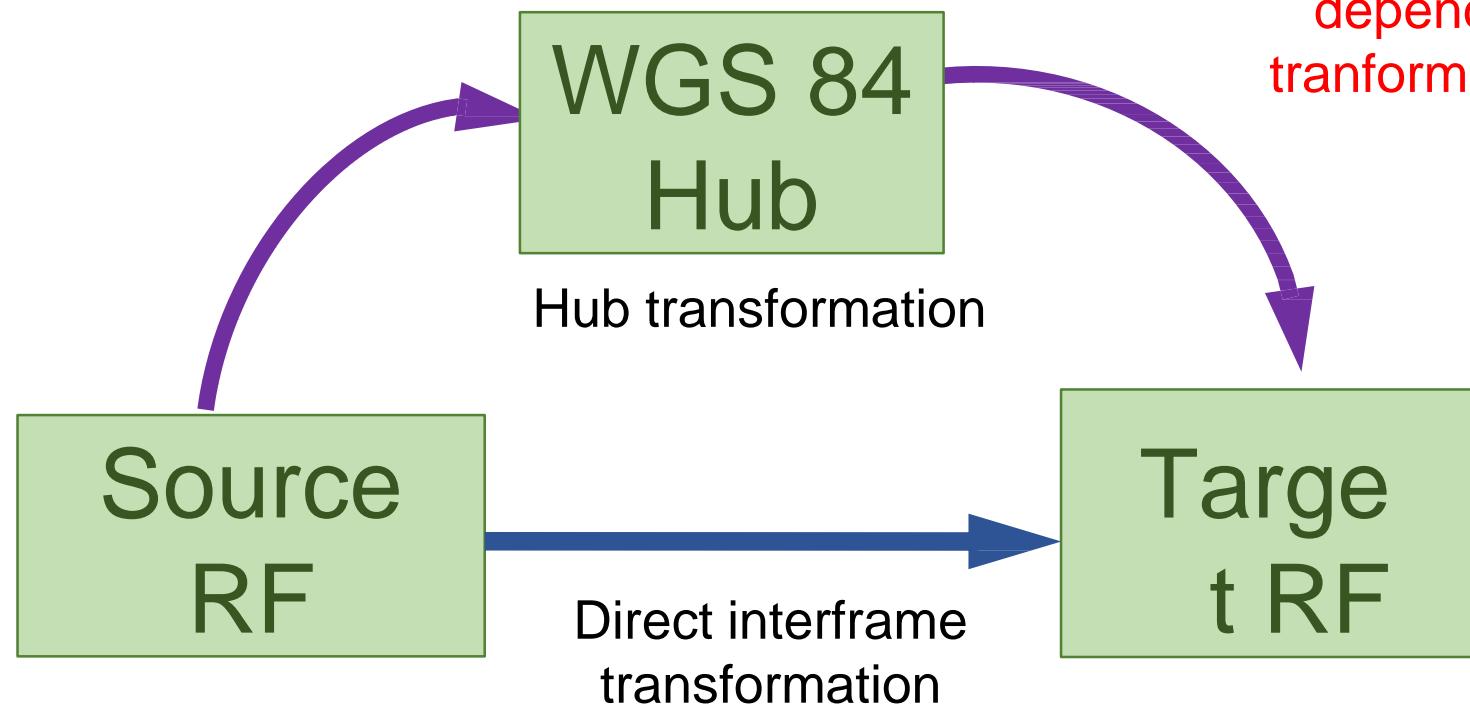
- All these transformations are used while working with GIS
- The main consideration is the alignment of precise data, defined in different RFs and **projections**.
- Geodetic Registries (EPSG and ISO Geodetic Registry) standardize definitions and transformation workflows for use in GIS, however there are still limitations with complex time-dependent RF.
- e.g. No standard for coseismic and postseismic displacement grids and no standardised epoch of WGS 84 when used as a hub transformation.
- Direct interframe transformations are the preferred option.

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## Reference Frames in Practice



### GIS transformations



# GIS Metadata Requirements

$\phi = 43^\circ 31' 32.3400'' \text{ S}$   
 $\lambda = 172^\circ 38' 23.4492'' \text{ E}$   
Geodetic NGZD2000 Epoch 2016.2

$\phi$  Precision:  $\pm 0.0001''$   
 $\lambda$  Precision:  $\pm 0.0002''$   
Epoch: 2016.2  
Type of position: Ellipsoidal  
Reference Frame: ITRF96  
Frame name: NGZD2000

X= -4593768.2707 m  
Y= 593377.9433 m  
Z= -4370031.2416 m  
Cartesian NGZD2000 Epoch 2016.2

X Precision:  $\pm 0.001 \text{ m}$   
Y Precision:  $\pm 0.002 \text{ m}$   
Z Precision:  $\pm 0.003 \text{ m}$   
Epoch: 2016.2  
Type of position: Cartesian  
Reference Frame: ITRF96  
Frame name: NGZD2000

$\phi, \lambda, h$

X, Y, Z

# GIS Metadata Requirements

$h = 23.126 \pm 0.007$  m Ellipsoidal NAD83(CSRS) Epoch 2013.2

Height: 91.256 m  
Precision:  $\pm 0.007$  m  
Epoch: 2013.2  
Type of height: Ellipsoidal  
Height system: NAD83  
Height frame name: CSRS

$h$

$H = 23.126 \pm 0.007$  m Orthometric CGVD2013(CGG2013) Epoch 2013.2

Height: 101.61 m  
Precision:  $\pm 0.01$  m  
Epoch: 2013.2  
Type of height: Orthometric  
Height system: CGVD2013  
Height frame: CGG2013

$H$

$N = -10.354 \pm 0.015$  m Geoidal CGG2013, NAD83(CSRS)

Geoid Height: -10.354 m  
Precision:  $\pm 0.015$  m  
Epoch: Static  
Type of height: Geoidal  
Model: CGG2013  
Frame: NAD83(CSRS)

$N$

$$H = h - N$$

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## Reference Frames in Practice



**Thank you!  
Gracias!**