IAG/FIG Commission 5/ICG Technical Seminar

 Reference Frame in Practice

 Rome, Italy 4–5 May 2012

Four Dimensional Deformation Models for Terrestrial Reference Frames

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- Introduction (10min Graeme)
- Concepts of 4d datums (10 min Graeme)
- The pros and cons of static, semi-dynamic datum and dynamic datums (10 min Grame)
- Development of Deformation Models (15min Richard)
- Incorporating the effects of events such as earthquakes into the model. (15min Richard)
- Case studies,
 - Australia (10 min Richard)
 - New Zealand, (10 min Graeme)
- Questions 10 min







Introduction





Fundamental Role of Reference Frame









Fundamental Role of Reference Frame









Fundamental Role of Reference Frame



Requirements of a National Reference System

- A coordinate framework that is accurate, stable, reliable and accessible
- Direct linkage to International Reference Frames
- **Simple** for users to connect to and use
- Physical infrastructure may include GNSS CORS and traditional geodetic survey marks
- Systems and tools to allow connection to the coordinate reference system and transformation of legacy data to the current reference system









Crustal Dynamics



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Temporal and spatial extent of deformation











Concepts of a 4D Datum









Static Datum (2D and 3D)

- Coordinates are fixed at a reference epoch
- Does not incorporate the effects of plate tectonics and deformation events Coordinates slowly go out of date, need to change periodically which is disruptive

Dynamic datum (4D)

Incorporates a deformation model to manage changes (plate tectonics and deformation events) Coordinates change continuously Can be confusing and difficult to manage

Semi - dynamic datum

Incorporates a deformation model to manage changes (plate tectonics and deformation events)

Coordinates fixed at a reference epoch

Change to coordinates is minimised







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The ideal world!

Need to accommodate error in model or changes in deformation Need to accommodate local and spatially complex deformation









Options - Error in model or changing deformation



Steer to a new model Jump to the new model Revise the previous model Ignore the previous model





Solution Revise the previous model always preserves the best estimate of past and future position and velocity





Options - Spatially complex deformation



The deformation event is not incorporated into the deformation model (dot is the base epoch coordinate of the mark)

The 'patch' deformation model – in this case a discrete event

The trajectory of the mark – incorporated the national deformation model and the 'patch'

The base epoch coordinate is changed to incorporate the offset calculated from the patch

Coordinates for times after the event just use the national deformation model and coordinates before the event include the patch.









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Episodic events



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Multiple events



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Earthquake and postseismic deformation



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Modelling the deformation



Options – regional deformation

- Simple rectangular grid (simplest method)
- Complex grid (eg curvilinear grid)

Options – complex deformation

- Densify the national deformation model (model becomes very complex)
- need a detailed triangulated grid becomes complex
- define a local 'patch' for the model (covers the area of the event with zero deformation at the boundaries)
- change coordinates











The pros and cons of static, semidynamic datum and dynamic datums





Advantages of a Dynamic or Semidynamic Datum



Maintains alignment with underlying global reference frames - ITRF

Lengthen the life of the datum

New observations can be integrated with old observations

Spatial accuracy of the geodetic network/datum is maintained or increased

Enables non-expert users to be isolated from the complexities of the dynamics (semi-dynamic datum only)

For practical purposes appears as a static datum (semi dynamic datum only)





Disadvantages of a Dynamic or Semi-dynamic Datum



Limited by the accuracy of the deformation model

Model can become complex over time to incorporate the effects of deformation events (e.g. earthquakes)

Coordinates need to be time tagged – cause confusion (dynamic datum only)

Most users do not know how to use a deformation model which is required to work with a dynamic datum

If using real time systems (CORS networks) need to use the deformation model to manage real time coordinates (semi-dynamic datums only)









Accommodate vertical deformation

- vertical deformation trends may be obscured by much larger localised episodic or cyclic events
- triangulated or other irregular grid probably required

Latency

- may be considerable time between a deformation event and the 'patch' being implemented
- for discrete events deformation may continue for some time requiring different versions of the patch

Extension Offshore

- how do you model deformation offshore?
- offshore may need to incorporate global model express velocities as global rotations

Changing Reference Epoch

 may ultimately need to change the reference epoch once coordinates become inconveniently different from true current positions (semi-dynamic datum only)

Joining adjacent jurisdictions/datum's







Development of a Deformation Model







Development of a Deformation Model





Aim of a Deformation Model



Dynamic (kinematic) – ITRF



Application:

GNSS data processing & analysis (e.g. PPP, RTK, NRTK, DGPS, Static post-processing) Large-scale deformation analysis GGOS

Semi-dynamic (kinematic) datum Fixed epoch of ITRF

Application:

All other spatial applications (e.g. cadastral, engineering, mapping, precision agriculture, mining, LiDar products) terrestrial surveying (e.g. TLS, total-station)





Classification of **Deformation**

Results in changes in coordinates of working frame



dam

m

dm

cm

mm

Magnitude of deformation yr⁻¹



Deformation is "invisible" in working frame - secular model Secular Non-secular Deformation Deformation Volcanisn Rigid Interseismic bstraction plate Water deformation motion Slow-Slip Event **Post glacial** Rebound month year day week decade century millenium Period



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hour

Deformation Postseismic

Deformation

minute

Smic

OSel



Developing a Secular Deformation Model







Secular Model format



aligned with current ITRF

no scale change / rate no translation / rate rotation of axes + rate

6 parameter transformation or Rx, Ry, Rz, + rates 3 parameter Euler propagation Ωx, Ω y, Ω z

aligned with earlier ITRF realisation or non ITRS ellipsoid

scale change / rate translation / rate rotation of axes + rate

14 parameter transformation S, Tx, Ty, Tz, Rx, Ry, Rz, + rates

Gridded secular deformation model ITRF site velocities on a 1° or 0.1° grid

Site velocity estimated by bilinear interpolation (as used in geoid or grid distortion modelling)





Kinematic ITRF to Semi-kinematic ITRF by Transformation



14 parameter transformation

$$\begin{bmatrix} X \\ Y \\ Z \end{bmatrix}_{t_0} = \begin{bmatrix} T_x + \dot{T}_x(t-t_0) \\ T_y + \dot{T}_y(t-t_0) \\ T_z + \dot{T}_z(t-t_0) \end{bmatrix} + \left\{ S + \dot{S}(t-t_0) \right\} \begin{bmatrix} 1 & \left\{ R_z + \dot{R}_z(t-t_0) \right\} & -\left\{ R_y + \dot{R}_y(t-t_0) \right\} \\ -\left\{ R_z + \dot{R}_z(t-t_0) \right\} & 1 & \left\{ R_x + \dot{R}_x(t-t_0) \right\} \\ \left\{ R_y + \dot{R}_y(t-t_0) \right\} & -\left\{ R_x + \dot{R}_x(t-t_0) \right\} & 1 \end{bmatrix}$$

6 parameter transformation (no translation or scale)

$$\begin{bmatrix} X \\ Y \\ Z \end{bmatrix}_{t_0} = \begin{bmatrix} 1 & \{R_z + \dot{R}_z(t - t_0)\} & -\{R_y + \dot{R}_y(t - t_0)\} \\ -\{R_z + \dot{R}_z(t - t_0)\} & 1 & \{R_x + \dot{R}_x(t - t_0)\} \\ \{R_y + \dot{R}_y(t - t_0)\} & -\{R_x + \dot{R}_x(t - t_0)\} & 1 \end{bmatrix} \cdot \begin{bmatrix} X \\ Y \\ Z \end{bmatrix}_{t_0}$$

$$t_0$$
is the reference epoch (years)tis the epoch of measurement (years) T_x, T_y, T_z Translation parameters (m) $\dot{T}_x \dot{T}_y \dot{T}_z$ R_x, R_y, R_z Rotation parameters (radians) $\dot{R}_x \dot{R}_y \dot{R}_z$ SScale (unitless) \dot{S}

rate of change (m/yr) rate of change (radians/yr) rate of change (per yr)





Kinematic ITRF to Semi-kinematic ITRF by Propagation



_ **_**

Site velocity from Euler pole definition

• V							• X	
Λ	$\left[\Omega_{V}Z - \Omega_{Z}Y \right]$		X		X		71	
• Y	$= \left \Omega_{Z}^{T} X - \Omega_{X}^{T} Z \right \cdot 1E-6$	\rightarrow	Y	=	Y	+	• Y	$\cdot (t_0 - t)$
·Z	$\left\lfloor \Omega_X Y - \Omega_Y X \right\rfloor$		$\lfloor Z \rfloor$	t_0	_ <i>Z</i> _	t	ż	2.00 M

 t_0 reference epoch of the semi-kinematic datum (in decimal years)tepoch of measurement (in decimal years) $(\Omega_X, \Omega_Y, \Omega_Z)$ Euler pole (Cartesian rotation format) $(X, Y, Z)t_0$ semi-kinematic coordinates computed at the reference epoch (m)(X, Y, Z)tkinematic ITRF coordinates at the measurement epoch (m) $(\dot{X}, \dot{Y}, \dot{Z})$ ITRF site velocity estimated from the Euler pole definition or interpolated from the secular deformation model (m/yr)





Gridded Deformation Models



Regular grid deformation model

- standard ASCII format
- (latitude, longitude, latitude rate, longitude rate, vertical rate)
- 1°, 0.25° or 0.1° grid size
- bilinear interpolation
- similar format to geoid model
 - planar assumption < 0.01 mm/yr error for 1° grid size
 - accommodates some localised deformation and strain (depending upon grid size)

Limitations of rigid plate and 14 parameter models

localised deformation distributed over model does not work where differential geodetic rates occur assumes rigid or uniformly deforming tectonic plate







Incorporating Episodic Events (e.g. earthquakes) into a Deformation Model





Why episodic events need to be modelled in





Localised deformation should result in coordinate changes to reflect visible reality





Typical time-series in a deforming zone









Time-series modelling



Separating seismic and secular (interseismic) deformation from time-series

Seismic patch is a sum of all non-secular (episodic) deformation between reference and measurement epoch







Nested model for deformation patch



Model Inputs –

InSAR

LiDar & High-res imagery

analysis of seismic data

Repeat GNSS obs of dense passive network (Strong argument for maintaining passive geodetic infrastructure)

Terrestrial surveys







Reverse prediction of position at reference epoch



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using velocity model

(Reversed)



Two modes of deformation in practice



GISB PRTU PARI 20 mm/yr Interseismic velocity 20 mm/yr NZGD2000 velocity model 200 mm Seismic Patch 2011.0

secular model (blue)

patch model (green)

existing model (orange)



Example



Rover (PRTU) ITRF2008 Epoch 2011.008

S 38° 48' 51.0946" E 177° 41' 52.3646"

The ITRF site velocity from interseismic velocity model : E -0.0108 m/yr N 0.0217 m/yr

The seismic patch model at epoch 2011.0 $\Delta E = 0.183$ m

∆N 0.008 m

NZGD2000 (estimated from model) S 38° 48' 51.1026" E 177° 41' 52.3620" NZGD2000 (tabulated) S 38° 48' 51.1021" E 177° 41' 52.3619"

 Tabulated – estimated:

 ∆E -0.002 m
 ∆N 0.014 m







Rigid plate case study (Australia)





Australian Plate Deformation





purple arrows – tectonic movement, green lines – baseline changes per year







Effect of Rigid plate rotation on GNSS baselines





Network at measurement epoch

Australian Plate rotates at ~0.63° / Ma

=

5 mm rotation of a 30 km GNSS baseline after only 15 years

e.g. holding rigid plate coordinates at an early epoch fixed for static processing or RTK at later epochs





Far-field deformation effects on rigid plate





Far-field deformation from great earthquakes around the Australian margin (e.g. Mw8.1 23 December 2004 Macquarie Island, from Watson *et. al*, 2010)





Intraplate deformation





Meckering, WA, 1968

Images courtesy Geoscience Australia



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Localised deformation





Surface creep

Sedimentary basins and non-bedrock sites can be subject to significant localised deformation







Australian Example - using AUSPOS





1 User Data

All antenna heights refer to the vertical distance from the Ground Mark to the Antenna Reference Point (ARP).

Station (s)	Submitted File	Antenna Type	Antenna Height (m)	Start Time	End Time
0015	0015241w.080	SOK_GSR2700IS NONE	0.910	2008/08/28 22:05:00	2008/08/29 08:30:30





Using AUSPOS



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		G12	306	301	306	301	306	301		PRN / # OF OBS
Astronomical Information		G13	272	250	272	250	272	250		PRN / # OF OBS
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www.ga.gov.au/bin/gps.pl

Wait > 3 days for IGS Rapid Orbit solution or > 14 days for IGS Final Orbit solution Select correct antenna model and antenna height to ARP

END OF HEADER

-2910.355

-1610.422

-3437.059

2338.379

3900.293

08 8 30 22 51 20.0000000 0 10G 2G10G12G15G18G21G24G29G30G31 24831460.15646 24831455.37546 133535311.44941 104052947.61741

22217869.10948 22217865.13347 119800792.08241 93350726.30141

22770826.22748 22770823.46946 122706607.26241 95614991.48041

25105337.14847 25105336.94545 131929440.60941 102802147.03941

24208741.73447 24208738.10945 127217788.34041 99130720.00841

-2267.813

-1254.879

-2678.230

1822.113

3039,188



Using AUSPOS





AUSPOS GPS Processing Report

April 24, 2012

This document is a report of the GPS data processing undertaken by the AUSPOS Online GPS Processing Service (version: AUSPOS 2.0). The AUSPOS Online GPS Processing Service uses International GNSS Service (IGS) products (final, rapid, ultra-rapid depending on availability) to compute precise coordinates in ITRF anywhere on Earth and GDA94 within Australia. The Service is designed to process only dual frequency GPS phase data.

Date	User Stations	Reference Stations	Orbit Type
2008/08/28 22:05:00	0015	ADE1 BEE2 BUR2 CEDU MOBS	IGS final
		PARK STR1 SYDN TID1	

4.1 Cartesian, ITRF2008

Station	X (m)	Y (m)	Z (m)	ITRF2008 @
0015	-4361680.131	3242720.715	-3326809.666	28/08/2008

5.1 Coordinate Precision - Geodetic, One Sigma

Station	σ East (m)	σ North (m)	σ Up (m)
0015	0.001	0.001	0.003

! Kinematic coordinates at epoch of observation





Example



3.1 Cartesian, GDA94

	Station	Х	(m)	Y	(m)		Z (m)	
	0015	-4361679	.624	3242720	.803	-332681	10.422	-
3.2 Geo	detic, GRS	80 Ellips	soid, (GDA94				
Station	Lat	tude	L	ongitude	Elli	psoidal	Derive	ed AHD
		(DMS)		(DMS)	He	ight(m)		(m)
0015	-31 38 33.	60805 1	43 22	14.83619		87.237	7	73.468
3.3 MG	A Grid, G	RS80 E	llipso	id, GDA	4 94			
Station	East	;	North	Zone	Ellip	soidal	Derive	d AHD
	(m)		(m)		Heig	ht (m)		(m)
0015	724826.751	. 649672	9.544	54		87.237	73	.468
	T_x	= -0.06386	$\mathfrak{F}(m)$				-	and a
	T_y	= 0.00023(a)	m)					
	T_z	= 0.04521(s	m)					
	S_c	= 1.1308e -	- 08					
	R_x	= 1.07834e	-07(r	radians)				
	R_y	= 9.49620e	-08(r	radians)				
	R_z	= 9.37467e	-08(r	radians)				
The above tra	nsformation pa	rameters are	e only y	valid for the	e epoch í	28/08/200	8.	-1









After a long day hopping across dry parameter space, a cold beer is near.

Thank you!







New Zealand Case Study





But we don't live on a stable planet







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Limitations with NZGD49



Regional distortions up to 5m present Built up in a piecemeal fashion Incompatible with global systems It is of limited spatial coverage It is static







NZGD2000



1998 – NZ introduced NZGD2000 (ref epoch 1 Jan 2000)

- geocentric origin
- aligned with the ITRS
- ITRF96 with epoch 2000.0 coordinates

NZGD2000 - semi-dynamic datum

 generalised motion of points modelled using a deformation model





NZGD2000



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- Tied to ITRF96
- Generalised motion of points modelled using a constant velocity deformation model
- Epoch 2000
 coordinates generated
 at 2000.0









Semi-dynamic datum

- deformation model calculated by holding fixed the Aust plate and uses the euler rotation parameters to bring in terms of ITRF96
- velocities provided in a rectangular grid
 (0.1 degree) for ease of computations
- current deformation model has horizontal constant velocities only
- generated using repeat surveys between 1992 and 1998
- enables propagation of
 coordinates and observations
 between reference epoch and
 observation epoch









What has gone well

- User acceptance
- Implementation of the deformation model in LINZ
- Maintaining the accuracy of datum

Issues

- Managing the deformation model
- Accuracy of deformation model versus CORS real time positions
- Managing the spatial alignment of the cadastral system
- Misalignment of readjusted historic geodetic control with new surveyed geodetic control







Future developments



- Updating the Deformation Model
- Vertical Deformation Model
- CORS Real Time Tools for Managing Coordinates
- Tie to the ITRF Going Fully Dynamic?





Darfield Earthquake









InSAR interference pattern as a result of the Darfield earthquake



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Each coloured fringe represents 1.5 cm of ground displacement in line-of-sight to the satellite

Incoherent regions indicate ground damage





Creating a patch – Canterbury earthquakes







Impact on geodetic control



Range is based on the distance from the centre of the fault rupture.

Maximum Range	Geodetic marks	Cadastral control	Total marks
(km)	(order 5 or better)	(order 6 or better)	
0-20	223	4816	56835
20-40	1269	49538	565892
40-60	3176	28632	387606
60-80	673	3681	143593
80-100	487	2182	103995
Total	5828	88849	1257921





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Residuals



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- The incorporation of a deformation model in NZGD2000 has enabled
 - the life of the datum to be lengthened
 - new observations to be integrated with old observations
 - the accuracy of the datum to be maintained
- But
 - how complex deformation events will be incorporated in the model have yet to be fully determined and resolved







Four dimensional deformation models for Terrestrial Reference Frames

